

## INSTITUTE OF GEOLOGICAL AND NUCLEAR SCIENCES LIMITED

Wairakei Research Centre  
State Highway 1, Wairakei  
Private Bag 2000  
Taupo  
New Zealand  
Phone +64-7-374 8211  
Fax +64-7-374 8199

Gracefield Research Centre  
69 Gracefield Road  
PO Box 30368  
Lower Hutt  
New Zealand  
Phone +64-4-570 1444  
Fax +64-4-569 0600

Rafter Laboratory  
30 Gracefield Road  
PO Box 31312  
Lower Hutt  
New Zealand  
Phone +64-4-570 4637  
Fax +64-4-570 4657

Dunedin Research Centre  
764 Cumberland Street  
Private Bag 1930  
Dunedin  
New Zealand  
Phone +64-3-477 4050  
Fax +64-3-477 5232

Client Report 43878D.10

Evaluation of Wells and  
Coppersmith (1994)  
Earthquake and Fault  
Relationships in the  
New Zealand Context

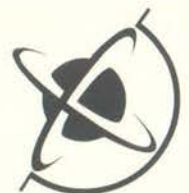
prepared for

Earthquake Commission  
Research Foundation  
EQC Project 97/249

by

Mark Stirling,  
David Rhoades, and  
Kelvin Berryman

December 1998



*Institute of*  
**GEOLOGICAL  
& NUCLEAR  
SCIENCES**  
*Limited*



*Institute of*  
**GEOLOGICAL  
& NUCLEAR  
SCIENCES**  
*Limited*

# **Evaluation of Wells and Coppersmith (1994) Earthquake and Fault Relationships in the New Zealand Context**

Mark Stirling, David Rhoades and Kelvin Berryman

Prepared for

Earthquake Commission Research Foundation  
EQC Project 97/249

Client Report 43878D.10

December 1998

COMMERCIAL-IN-CONFIDENCE

This report has been prepared by the Institute of Geological & Nuclear Sciences Limited exclusively for and under contract to The Earthquake Commission Research Foundation. Unless otherwise agreed in writing, all liability of the Institute to any party other than the The Earthquake Commission Research Foundation in respect of the report is expressly excluded.



## Contents

Contents page	i
Abstract	1
Introduction	1
Method and Analysis	2
Discussion and Conclusions	16
References	17
Appendix	18



### Abstract

We assess the suitability of the earthquake and fault relationships of Wells and Coppersmith (1994) for use in New Zealand seismic hazard studies. We find that the regressions of Wells and Coppersmith (hereafter referred to as "W&C") provide underestimates of the moment magnitudes ( $M_w$ ) and coseismic surface displacements of large New Zealand earthquakes, and attribute much of the discrepancy to the dataset used by W&C for their analysis. Regressions developed after addition of newly published data to the W&C dataset, and addition of data originally excluded from the W&C dataset provide closer estimates of  $M_w$  and displacement to those of the New Zealand earthquakes. The remaining discrepancies appear to be due to the New Zealand earthquake dataset comprising largely paleoearthquake and pre-instrumental data, which produce larger estimates of  $M_w$  and displacement than more recent data. Lastly, we find that the coseismic displacement per unit rupture length (proportional to stress drop) is not constant for all earthquakes, but is a decreasing function of both slip rate and the total amount of slip registered across the fault.

### Introduction

W&C (1994) developed a series of empirical regressions between earthquake magnitude and various fault rupture parameters from a worldwide dataset of historical earthquakes. These empirical regressions have since become the standard for use in seismic hazard analysis throughout the world. However, the regressions of W&C tend to provide underestimates of the magnitudes and single event displacements for large New Zealand earthquakes that have been observed historically or estimated from paleoseismic data (e.g. Stirling et al. 1998). For example, the M8.1-8.2 1855 Wairarapa earthquake, and M7-7.3 1888 Canterbury earthquake were about 0.5 and 0.2 magnitude units larger, respectively than the magnitudes estimated from the rupture lengths of those earthquakes with the regression equations of W&C. It is therefore possible that the use of W&C's regressions for seismic hazard analysis in New Zealand may result in underestimates in seismic hazard for large earthquakes. Another unresolved issue concerning the regressions of W&C is whether or not factors such as tectonic setting influence relationships between magnitude, displacement and other fault parameters. While W&C found no evidence for this in their analysis, the recent regression of Anderson et al. (1996) clearly shows that larger earthquakes tend to occur on slower slipping faults, for a constant rupture length.

In this study we aim to resolve the discrepancies between the predictions of W&C and the New Zealand data. Our general approach is to determine whether or not the discrepancies are due to



uniqueness in the New Zealand tectonic environment, or are an artifact of the data and methods used by W&C to construct their regression equations. W&C constructed regression equations from the source parameters of  $M_w \geq 4.7$  earthquakes, but excluded a lot of earthquakes from their analysis that they considered too unreliable to use. Their largest exclusions involved almost all early (approximately pre-1940) earthquakes, including several New Zealand earthquakes. Their exclusions therefore involved early instrumental data (the first seismograms date from the late 19<sup>th</sup> century), and pre-instrumental data derived from felt intensities and from field measurements of historical ruptures. In this study we compile a large dataset that includes much of the early data W&C excluded from their analysis, and also include newly published data for early historical (pre-instrumental) to recent earthquakes from the eastern Mediterranean region (Ambraseys and Jackson, 1998), pre-instrumental data from Japan (Research Group for Active Faults of Japan, 1991), and a mixture of pre-instrumental and recent data from the regressions of Anderson et al. (1996). We then construct regression equations from our worldwide dataset, and equations based on New Zealand data alone, and compare our regressions to W&C's original regressions. Since there are relatively few New Zealand earthquakes in W&C's original dataset, we boost the New Zealand component of the dataset by adding estimates of  $M_w$  and fault rupture parameters from paleoearthquake data for New Zealand active faults. We also add paleoearthquake data from Japan to increase the size of the global dataset. Our dataset therefore has considerably wider scope than any of the previous datasets used to produce empirical regressions of earthquake magnitude from fault parameters.

### Method and Analysis

The estimates of earthquake magnitude and fault parameters used in this study are listed in the Appendix. In all, 283 earthquakes are listed in the dataset. Almost all of the time spent on this project went into evaluating and compiling data from existing regression analyses, searching the literature for data, and obtaining data by communication with colleagues in New Zealand and overseas. Each row in the Appendix represents a single earthquake, and provides data on the earthquake date, location, magnitude, slip type, rupture length (surface and subsurface estimates), rupture width, single event displacement (maximum and average values of surface displacement), and the values of slip rate and total slip for the associated fault. A detailed description of the dataset is given at the base of the Appendix. Since more than one estimate of magnitude is available for many of the earthquakes, we select the best estimate of magnitude from the following choices, in decreasing order of preference; (i)  $M_w$  from seismological data, (ii)  $M_s$  or  $M_{jma}$  (surface wave magnitudes) as an approximation of  $M_w$  (iii) MI (intensity



magnitude) as an approximation of  $M_w$  and (iv)  $M_w$  from paleoearthquake data. We restrict our dataset to those earthquakes with  $M_w \geq 6.0$ , which is the approximate lower limit of  $M_w$  that is commonly associated with surface rupture (e.g. W&C 1994). W&C incorporated some earthquakes of  $M_w < 6.0$  into their analysis, but we have verified that these events did not influence their regressions significantly. However, we include in our analysis all of the pre-1940 data listed by W&C but not used in their analysis.

We develop regression equations of a similar form to those developed by W&C. The first form is used to relate  $M_w$  to the rupture length or area:

$$M_w = a + b(\log(L \text{ or } A)),$$

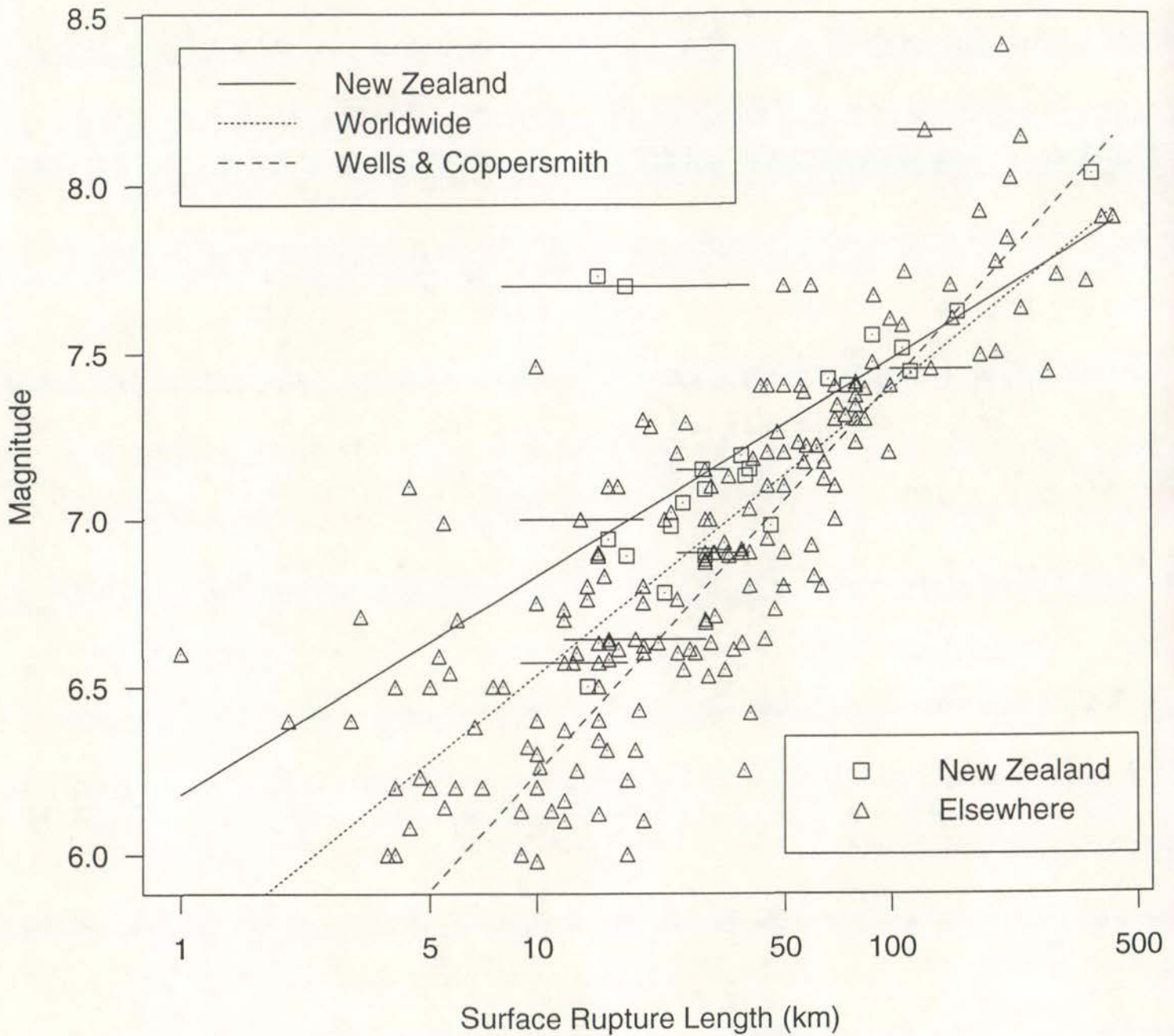
in which  $L$  is the surface rupture length in km,  $A$  is the fault area in  $\text{km}^2$ ,  $a$  and  $b$  are empirically-derived constants, and  $\log$  is to the base 10. The second form of equation is used to relate the average value of single event displacement to the rupture length:

$$\log(D) = a + b(\log(L))$$

in which  $L$  is the surface rupture length in km, and  $D$  is the average value of coseismic displacement (in metres) measured at the ground surface along  $L$ . We follow W&C's methodology and construct regressions of magnitude versus surface rupture length, and displacement versus surface rupture length. The distinction between surface and subsurface (aftershock) derived estimates of rupture length is important. Surface rupture length tends to be less than subsurface rupture length for a given earthquake, leading to systematic differences in the regressions of earthquake magnitude versus rupture length. We also construct regressions of  $M_w$  versus rupture area, in which the rupture area is calculated from the estimates of surface rupture length and downdip width of the rupture. Our emphasis on surface rupture parameters in the above regressions is driven by a need to more correctly estimate the magnitudes of future large New Zealand earthquakes from the length of surface ruptures, and from the length of fault traces in general. We therefore do not present any regressions based on subsurface estimate of rupture length in this study.

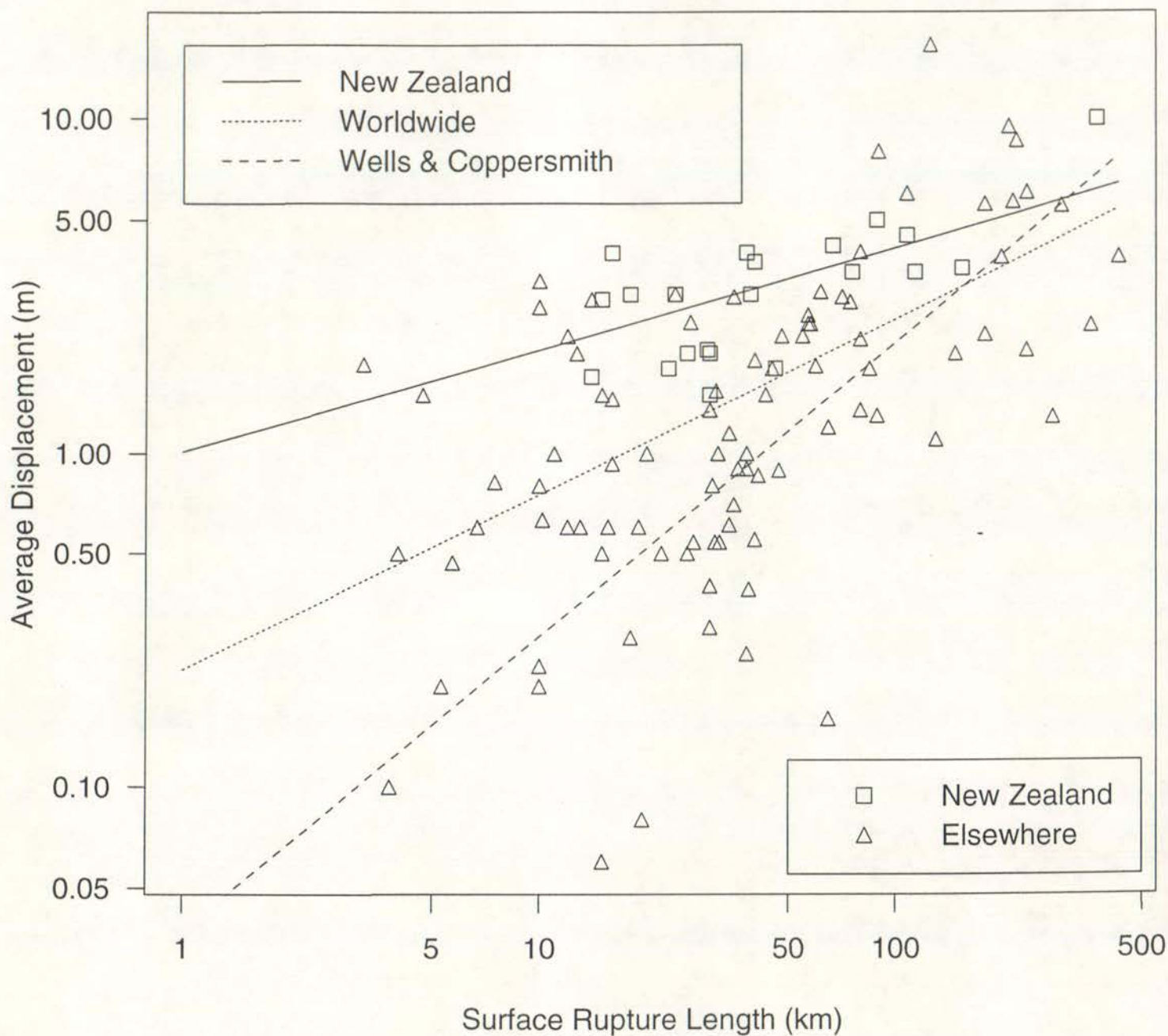
We show the results of our regression analyses of  $M_w$  versus surface rupture length, surface rupture displacement versus surface rupture length, and  $M_w$  versus rupture area in Figures 1 to





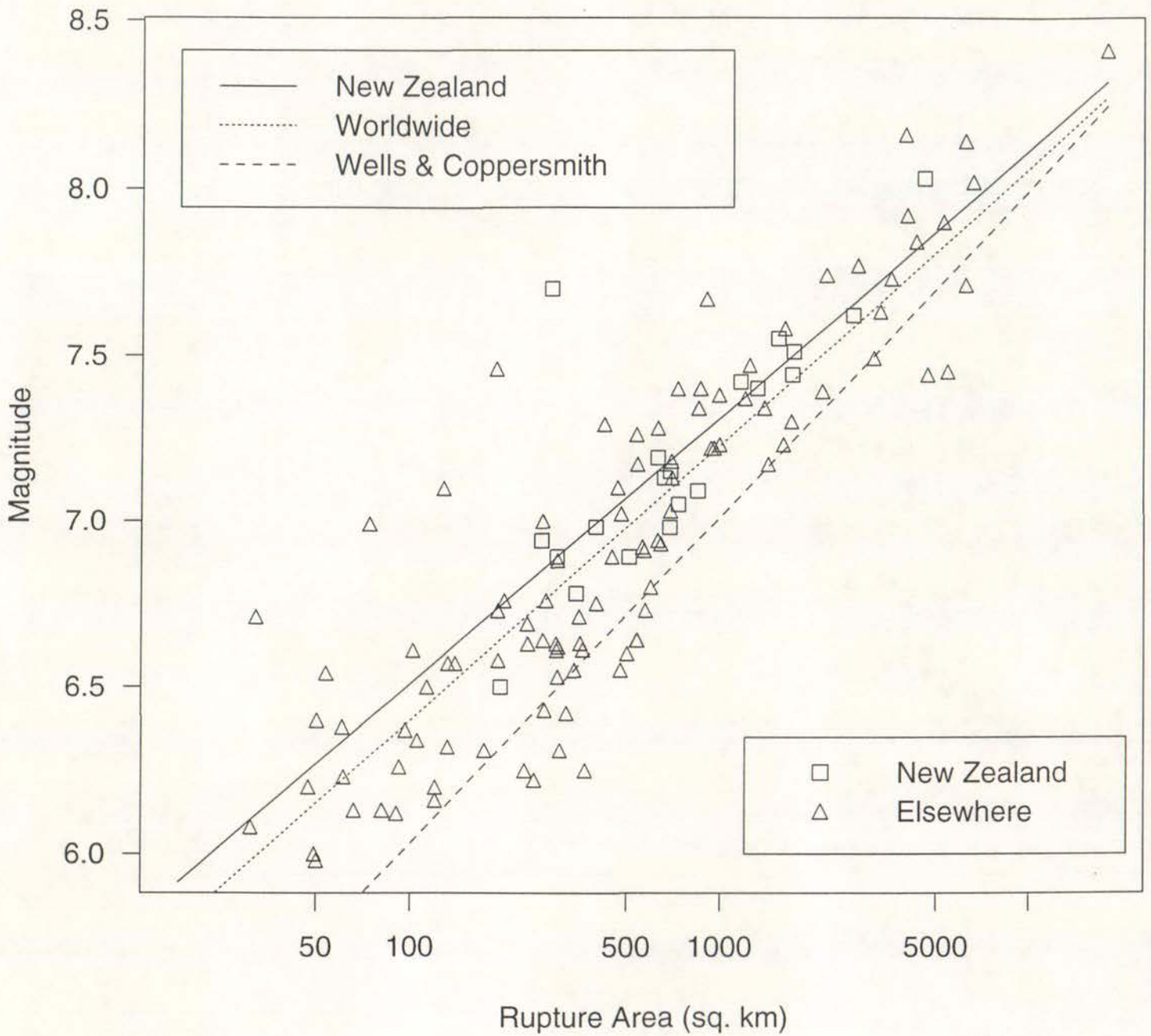
**Figure 1a**

Regressions of magnitude versus surface rupture length for our worldwide dataset, the New Zealand subset of the dataset, and the original regressions of W&C (labelled "Wells and Coppersmith"). The error bars on some symbols reflect uncertainties in estimates of rupture length for a given earthquake. See Table 1 for parameters  $a$  and  $b$  for each of the regressions.



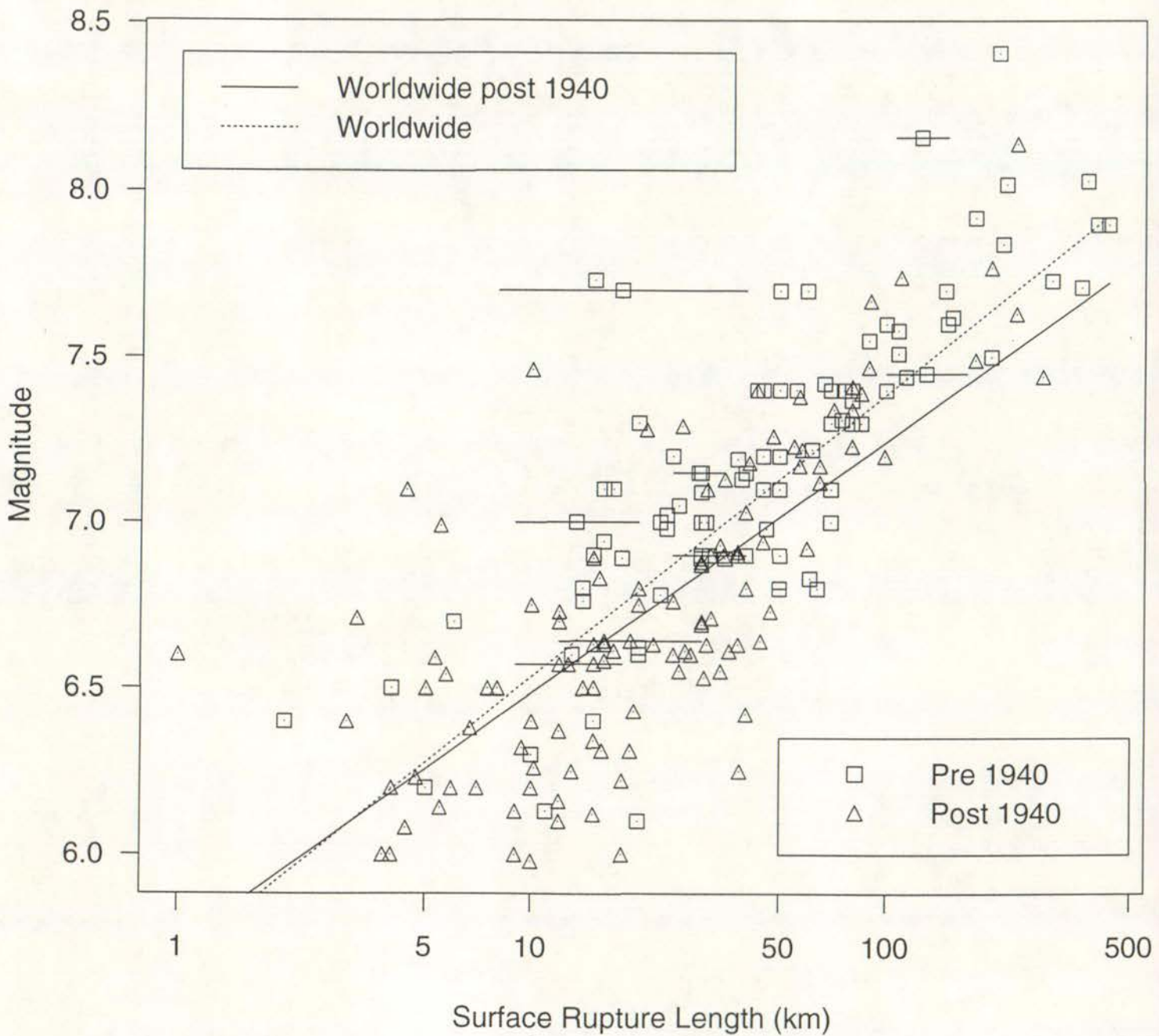
**Figure 1b**

Regressions of average surface rupture displacement versus surface rupture length for our worldwide dataset, the New Zealand subset of the dataset, and the original regressions of W&C (labelled "Wells and Coppersmith"). See Table 1 for parameters  $a$  and  $b$  for each of the regressions



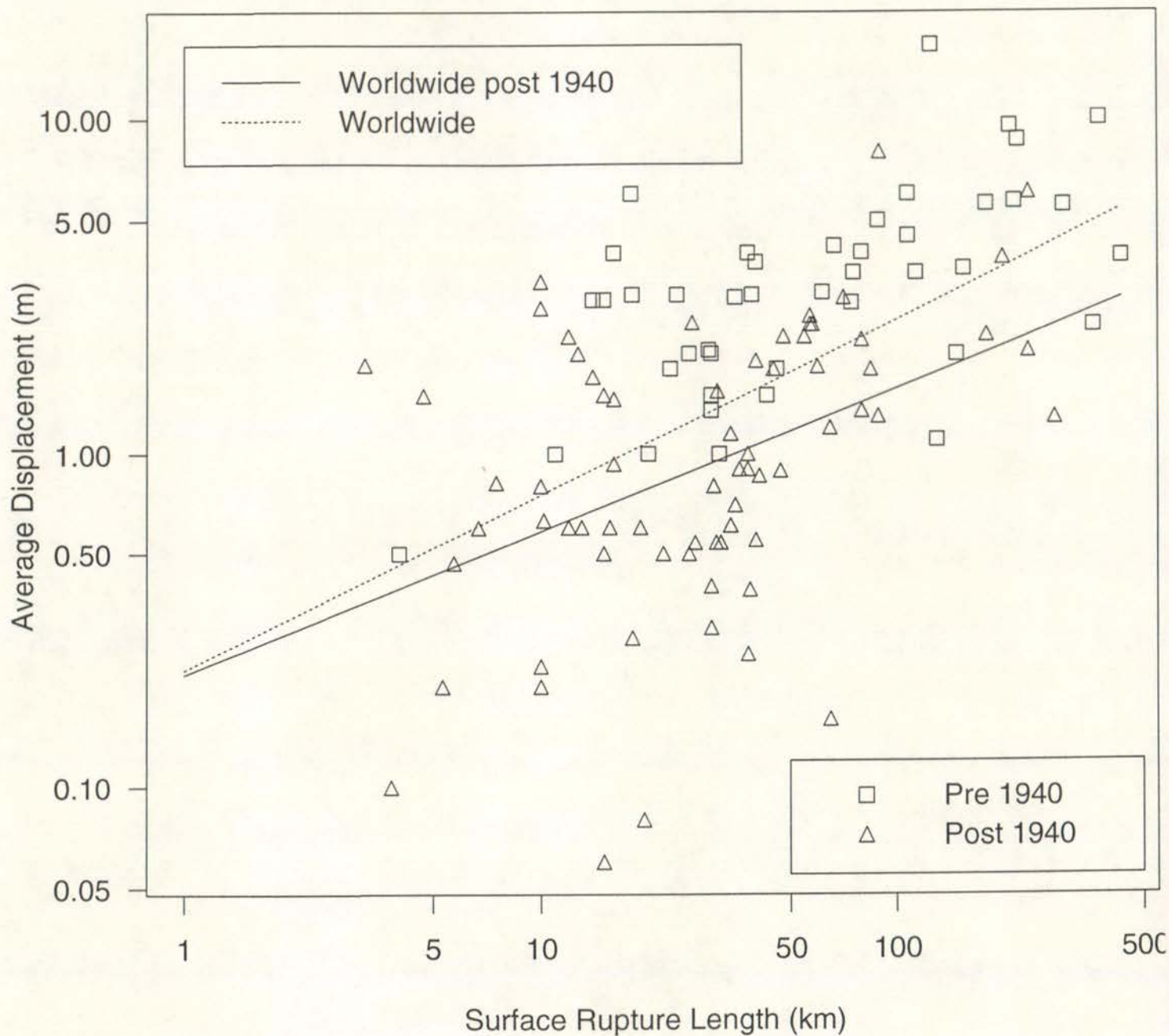
**Figure 2**

Regressions of magnitude versus rupture area, in which rupture area is the product of the surface rupture length and the downdip width of the rupture. See Table 1 for parameters  $a$  and  $b$  of the regressions.



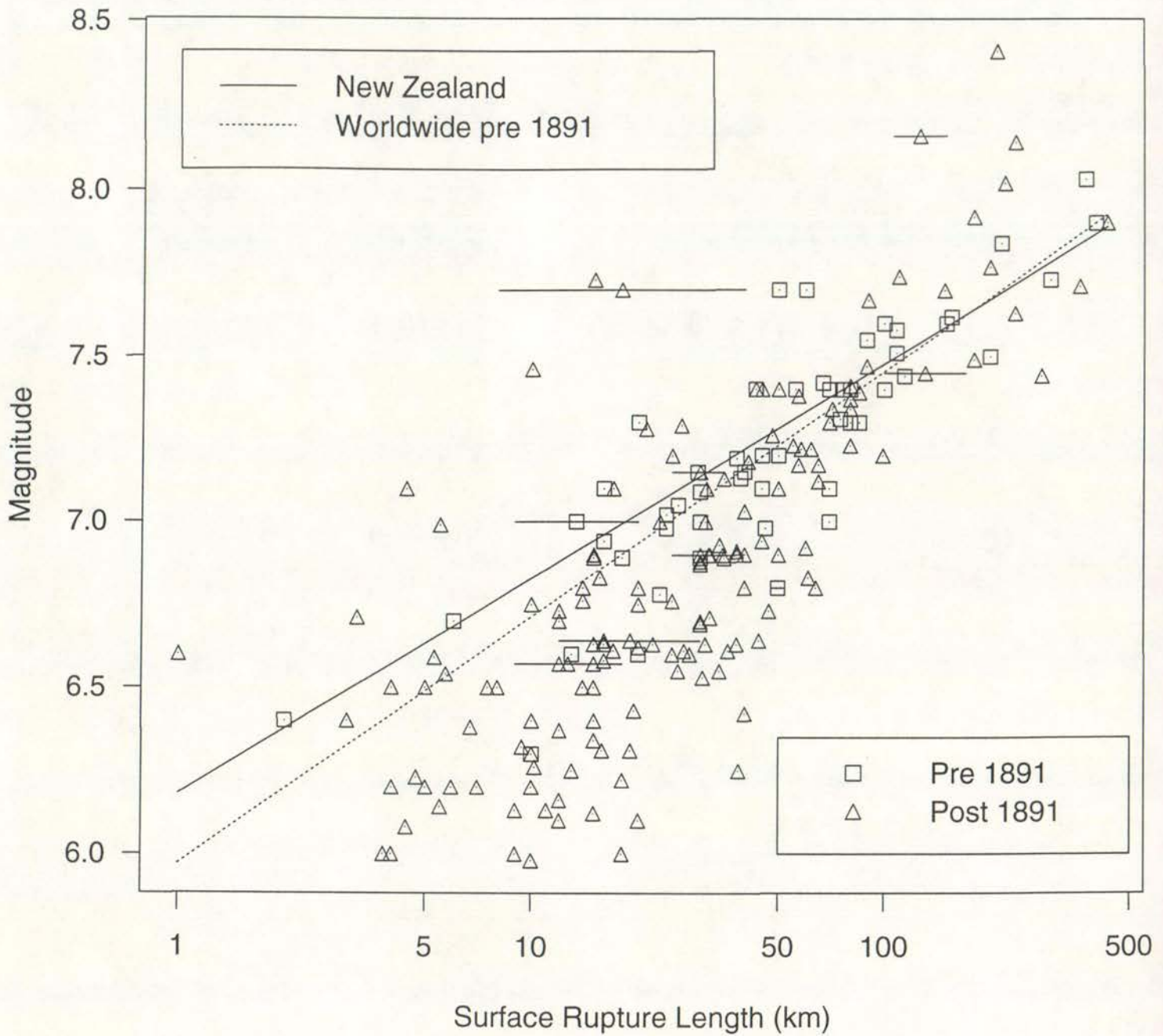
**Figure 3a**

Regressions of magnitude versus surface rupture length for our worldwide dataset, and for data from 1940 onwards). The error bars on some symbols reflect uncertainties in estimates of rupture length for a given earthquake. See Table 1 for parameters  $a$  and  $b$  of the regressions.



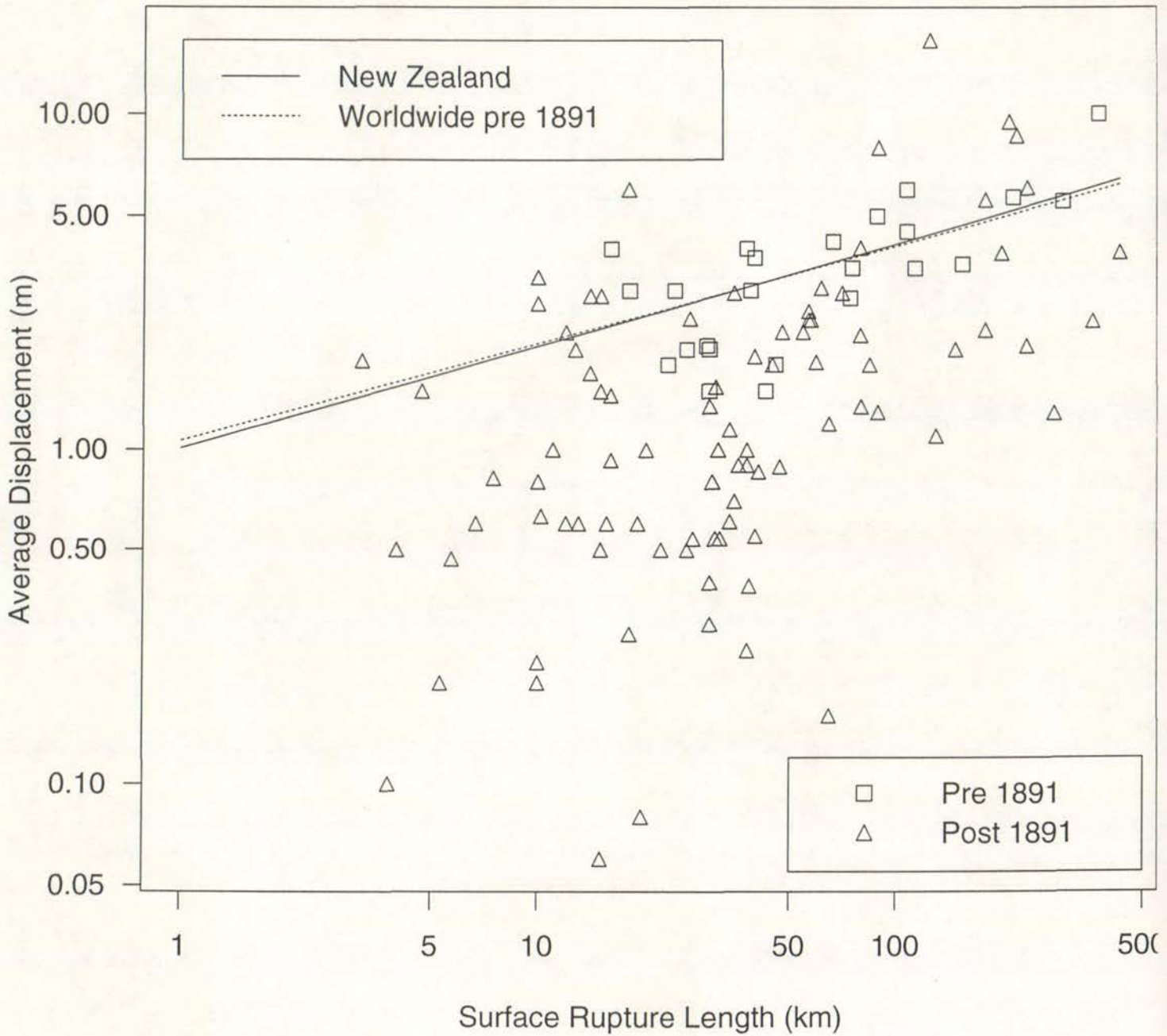
**Figure 3b**

Regressions of average surface rupture versus surface rupture length for our worldwide dataset, and for worldwide data from 1940 onwards. See Table 1 for parameters a and b of the regressions.



**Figure 4a**

Regressions of magnitude versus surface rupture length for the New Zealand subset of our dataset, and for all pre-instrumental (pre 1891) data). The error bars on some symbols reflect uncertainties in estimates of rupture length for a given earthquake. See Table 1 for parameters  $a$  and  $b$  of the regressions.



**Figure 4b**

Regressions of average surface rupture displacement versus surface rupture length, for the New Zealand subset of our dataset, and for all pre-instrumental (pre 1891) data. See Table 1 for parameters a and b of the regressions.



**Table 1: Regression Parameters**

**1. Regressions of Mw on surface rupture length (L)**

$$M_w = a + b \times \log(L)$$

Subset	N	a(sa)	b(sb)	R.s.d.
All	206	5.68(0.07)	0.85(0.04)	0.309
1940 and later	113	5.73(0.09)	0.76(0.06)	0.310
Pre 1891	54	5.97(0.10)	0.74(0.06)	0.186
NZ	23	6.18(0.27)	0.65(0.16)	0.278
W&C All (original)	77	5.08(0.10)	1.16(0.07)	0.28

**2. Regressions of Mw on fault area (A)**

$$M_w = a + b \times \log(A)$$

Subset	N	a(sa)	b(sb)	R.s.d.
All	122	4.75(0.11)	0.83(0.04)	0.262
NZ	21	4.91(0.40)	0.80(0.14)	0.224
W&C All (original)*	148	4.07(0.06)	0.98(0.03)	0.24

\* Rupture areas are based on the spatial extent of aftershocks in W&C's original regressions.

**3. Regressions of average surface displacement (D) on surface rupture length (L)**

$$\log(D) = a + b \times \log(L)$$

Subset	N	a(sa)	b(sb)	R.s.d.
All	116	-0.65(0.13)	0.52(0.08)	0.392
1940 and later	70	-0.66(0.16)	0.43(0.11)	0.385
Pre 1891	27	0.03(0.16)	0.29(0.09)	0.182
NZ	23	0.01(0.16)	0.30(0.10)	0.166
W&C All (original)	66	-1.43(0.18)	0.88(0.11)	0.36

**4. Regressions of the ratio displacement/surface rupture length (D/L) on slip rate (SR) or total slip (TS)**

$$\log(D/L) = a + b \times \log(SR)$$

Subset	N	a(sa)	b(sb)	R.s.d.
All	63	-1.28(0.05)	-0.30(0.06)	0.349

$$\log(D/L) = a + b \times \log(TS)$$

Subset	N	a(sa)	b(sb)	R.s.d.
All	27	-1.54(0.07)	-0.15(0.08)	0.337

N=Number of data used in regression, a and b are the parameters a and b of the regression (with associated standard error in parentheses), R.s.d is the standard deviation for the dependant variable.





4, and provide a table of regression parameters in Table 1. For comparison, Table 2 is constructed in a similar manner to the table of parameters given by W&C (1994). We show the regressions of  $M_w$  and displacement versus surface rupture length for our worldwide dataset, the New Zealand component of the dataset, and the original regressions of W&C in Figure 1. The most obvious trends in Figure 1 are that our New Zealand regressions show larger  $M_w$  and displacements than our worldwide regressions at rupture lengths less than 200km, and that both our New Zealand and worldwide regressions show larger  $M_w$  and displacements than the regressions of W&C, for the same range of rupture lengths. Specifically, our regression lines have smaller slopes (parameter  $b$ ) and larger  $y$  intercepts (parameter  $a$ ) than W&C's regression lines. Similar discrepancies are observed for regressions of  $M_w$  versus rupture area (Fig. 2). The differences between our worldwide regressions and those of W&C must somehow arise from the incorporation of the new data into our analysis. Since a large proportion of these new data are pre-1940 in age (i.e. not used by W&C because they were considered too unreliable for regression analysis), it is possible that they introduce a systematic bias to our regressions. Our new regressions do not appear to simply be the by-product of a larger random scatter in the data as compared to the dataset of W&C, since the standard deviations for  $M_w$  from our regressions and those of W&C are similar ("R.s.d" in Table 1). Our approach to determining whether the pre-1940 data have introduced a systematic bias to our worldwide regressions is to compare regressions constructed from the younger (1940 onwards) subset of our dataset to our worldwide regressions in Figure 3 to examine the influence of the pre-1940 data on the latter. Since the regression curves for the 1940 onwards data are very similar to those of our worldwide curves in Figure 3 it is clear that the pre-1940 earthquake data do not introduce a bias to our worldwide curve. The differences between our worldwide regression and that of W&C (Fig. 1) are therefore due to the specific choice of data in W&C's regressions. As well as excluding almost all pre-1940 data, they also excluded a number of the more recent earthquakes from their analysis. On examination of graphs in W&C and comparison with our Figure 1 it appears that they removed all of the "outlier" data, meaning data that would increase the overall scatter in their graphs, and possibly increase the uncertainties in their analysis if included.

Increasing the size of our global database over that used by W&C accounts for about half of the discrepancy between our New Zealand regression curve and the original regression of W&C (Fig. 1). The remaining discrepancy in Figure 1 may therefore be due to unique scaling relationships between  $M_w$  or displacement and rupture length in the New Zealand environment. However, it is difficult to envisage why this would be so, since the New Zealand earthquake data come from a great variety of tectonic regimes. The discrepancies could alternatively be due

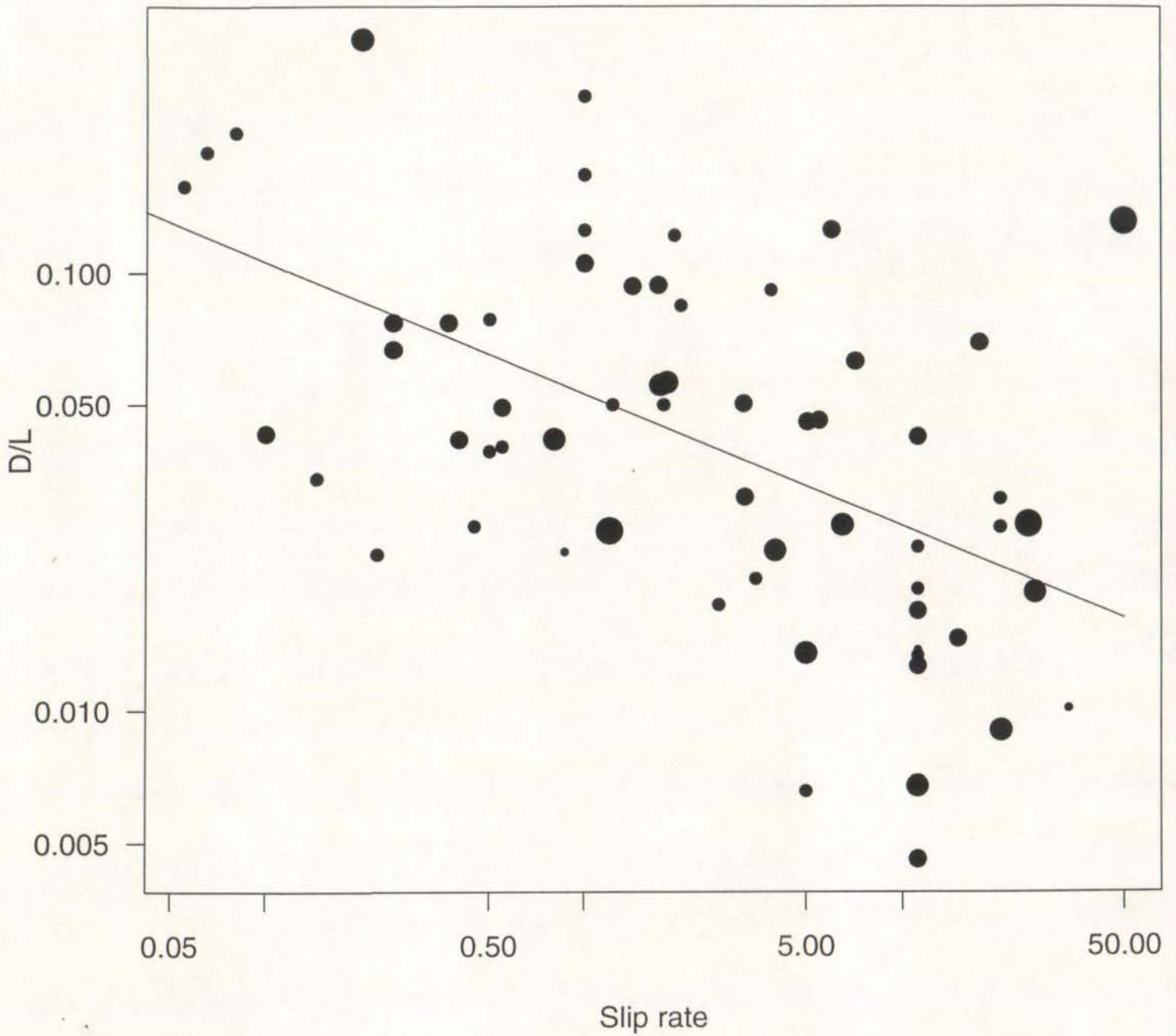


to an upward bias in the estimates of  $M_w$  and displacement from pre-instrumental (i.e. pre 1891, the year of the earliest instrumentally recorded earthquake in the Appendix) and paleoseismic data, since over 80% of the New Zealand earthquakes listed in the Appendix are from these sources. To determine which of these explanations is the more realistic, we construct a regression for all of the pre-instrumental (pre-1891 and paleoseismic) data in the Appendix, and compare these to our New Zealand regressions in Figure 4. Since the regression curves all plot very close together in Figure 4, it appears that the remaining discrepancies between our New Zealand and global regressions in Figure 1 are due to the dominance of pre-instrumental and paleoearthquake data in the New Zealand subset of the data. Since these data only form a small proportion of our worldwide dataset as a whole, they have not produced a noticeable bias in the worldwide regressions.

Systematic overestimation of  $M_w$  from paleoseismic and pre-instrumental data could arise from underestimation of surface rupture length and/or overestimation of the average value of coseismic displacement along the rupture. Since the amount of coseismic displacement along a rupture tapers to zero at the ends of a rupture, scarp degradation processes would progressively obliterate the ends of a rupture, leading to reduction of the observable rupture length over time. It is also conceivable that the processes of scarp degradation would tend to preferentially remove the smallest displacements along a fault scarp, which could lead to overestimation of the average values of displacement along the fault. The discrepancies in Figure 1 may also be due to our New Zealand regressions being poorly constrained at rupture lengths less than about 20 km, in contrast to the wide range of rupture lengths represented by the global dataset. Again, this could be attributed to scarp degradation processes, which would lead to obliteration of the short rupture lengths and associated small displacements. In general the smallest displacements and rupture lengths would be more difficult to identify in the field than the larger rupture lengths and displacements, regardless of their state of preservation.

Lastly, we examine the influence of slip rate and total slip on stress drop, by constructing regressions between the displacement/rupture length (equivalent to stress drop) and the total slip, and displacement/length and the slip rate from our global dataset in the Appendix. The form of the two equations is:

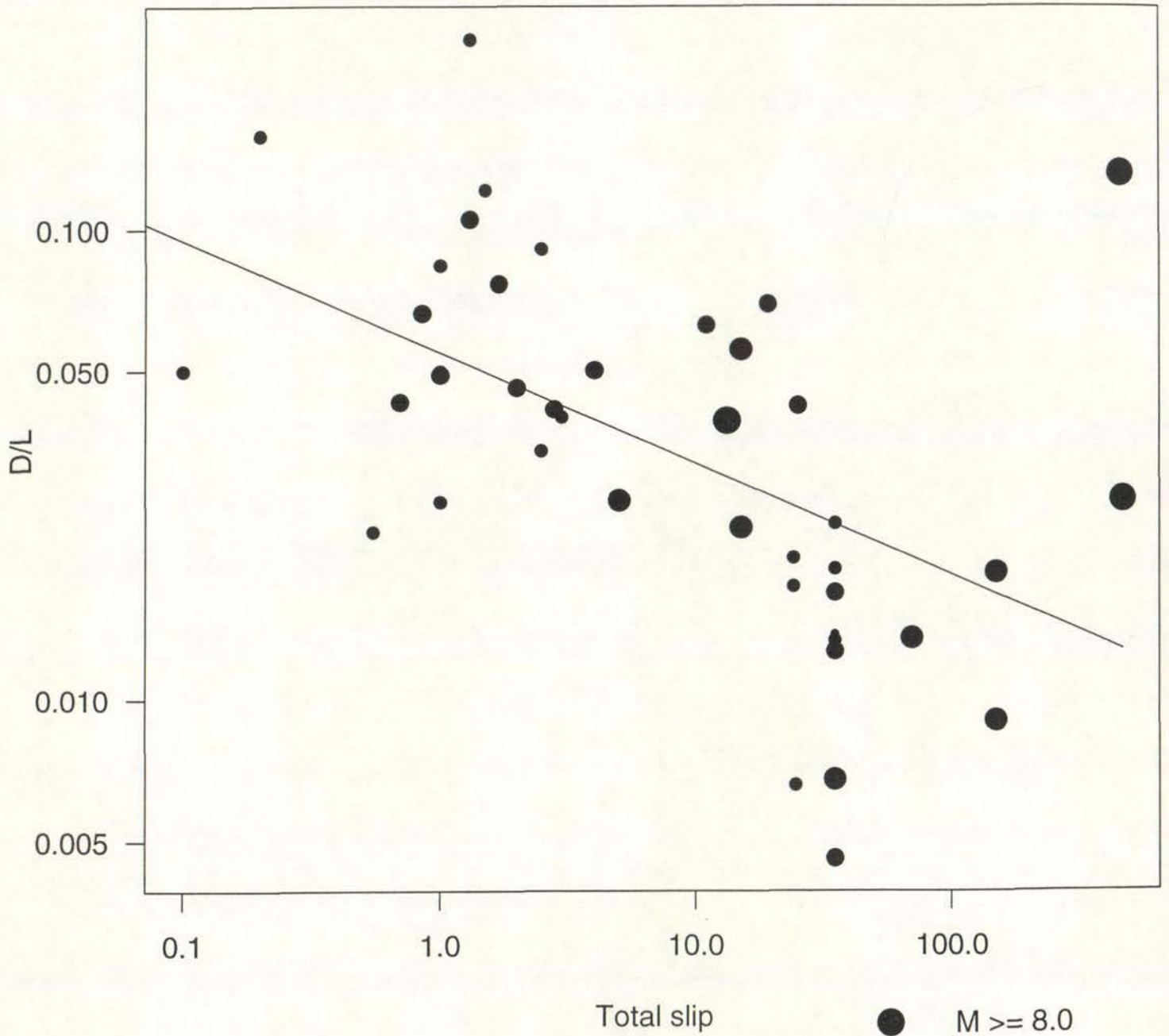
$$\text{Log}(D/L) = a + b(\text{log}(SR \text{ or } TS)),$$



- $M \geq 8.0$
- $7.5 \leq M < 8.0$
- $7.0 \leq M < 7.5$
- $6.5 \leq M < 7.0$
- $6.0 \leq M < 6.5$

**Figure 5a**

Plots of displacement per unit rupture length versus fault slip rate. The size of the symbols are proportional to the  $M_w$  of the earthquake.



**Figure 5b**

Plots of displacement per unit rupture length versus the total slip registered across the causative fault. The size of the symbols are proportional to the  $M_w$  of the earthquake.



in which  $D$  is the average displacement (m),  $L$  is the rupture length (km),  $SR$  is the preferred value of slip rate, (mm/yr), and  $TS$  is the value of total slip recorded across the fault (km). The parameters  $a$  and  $b$  for each of these regressions are given in Table 1, and the regression lines are shown in Figure 5. The radius of each data point is proportional to  $M_w$  on the two graphs.

Both of the graphs in Figure 5 show considerable scatter. However, there appears to be a tendency for the surface displacement/rupture length to be a decreasing function of fault slip rate and total slip. These results suggest that stress drop is not constant for all earthquakes, but can vary by more than an order of magnitude depending on both the present and long term activity of the associated fault.

### Discussion and Conclusions

Our analysis shows that the magnitudes and associated coseismic displacements for large New Zealand earthquakes are generally larger than the values predicted from the regressions of W&C for the same rupture length. The discrepancies are a decreasing function of rupture length. We attribute the discrepancies to the dataset used in our study versus that of W&C. Regressions developed after addition of newly published data to the W&C dataset, and addition of data originally excluded from the W&C dataset (almost all pre-1940 data, along with some later data) provide closer estimates of  $M_w$  and displacement to those of the New Zealand earthquakes. We are unsure of the exact reasons that W&C made their data selection, since the uncertainty in their regression-derived estimates of  $M_w$  (as reflected by the standard deviations given in Table 1) are basically no less than ours. We also find that pre-instrumental (i.e. pre-1891) and paleoearthquake data tend to provide overestimates of  $M_w$  and coseismic displacement, and the dominance of these data in the New Zealand subset of our dataset have resulted in regressions for New Zealand that show large estimates of  $M_w$  and displacement for a given rupture length. While the standard deviations given for each of the regressions in Table 1 are such that the observed differences in Figures 2 to 4 are rarely significant beyond the 1 sigma level, the observed differences between the regression curves translate to large differences in mean estimates of  $M_w$  or displacement for a given surface rupture length. These large differences in mean estimates of  $M_w$  or displacement from regression equations in general should be thoroughly addressed in the treatment of parameter uncertainty in seismic hazard analysis. In particular, the large uncertainties in estimates of  $M_w$  and coseismic displacement that arise from the use of New Zealand pre-instrumental and paleoseismic data that can be inferred from Figures 1 to 4 should also be taken into account in seismic hazard studies, and efforts focused on reducing these uncertainties. In general, pre-instrumental (pre-1891) and paleoseismic data



should be excluded from future regression analyses if they form the majority of a regression dataset. A formal assessment of the uncertainties in these data should be conducted. However, we see no problems with the general use of post 1891 data for regression analysis.

Lastly, our results indicate that stress drop (proportional to coseismic displacement per unit rupture length; Fig. 5) is not constant for all earthquakes, and that a link exists between the stress drop of earthquakes and the physical features of a fault. Specifically, the highest stress drops appear to be associated with the faults that show the smallest amounts of cumulative slip, and slowest slip rate. If stress drop correlates with the amplitude of strong ground motions (e.g. Anderson et al., 1996) then we might expect the strongest ground motions to occur on faults with the smallest amounts of cumulative slip and lowest slip rates. Incorporation of parameters such as slip rate and total slip into regressions may therefore provide useful information on the ground motions that are likely to accompany the occurrence of large earthquakes on a particular fault.

### References

- Aki, K. and Richards, P.G. (1980): *Quantitative seismology: theory and methods*. San Francisco, California. W.H. Freeman.
- Ambraseys, N.N. and Jackson, J.A. (1998): Faulting associated with historical and recent earthquakes in the Eastern Mediterranean region. *Geophysical Journal International* 133: 390-406.
- Anderson, J.G., Wesnousky, S.G., and Stirling, M.W. (1996): Earthquake size as a function of fault slip rate. *Bulletin of the Seismological Society of America* 86: 683-690.
- Research Group for Active Faults of Japan (1991): *Maps of active faults in Japan with an explanatory text*. University of Tokyo Press.
- Stirling, M.W., Wesnousky, S.G., and Berryman, K.R. (1998). Probabilistic seismic hazard analysis of New Zealand. *New Zealand Journal of Geology and Geophysics* 41 (in press).
- Wells, D. and Coppersmith, K. (1994): New empirical relationships among magnitude, rupture length, rupture width, rupture area, and surface displacement. *Bulletin of the Seismological Society of America* 84: 974-1002.











USA	NA	Northridge	17/1/94	R		6.7	6.7		8	16		17.5	23.3		1.2		1	1	0.3	0.3	0.3	148	5 x	20,24
R	MT	Lut	23/2/94	FS	6		6				4	4												
NZ	NZ	Arthurs Pass	18/6/94	R		6.7	6.7		30	40			4	7										23
GR	MT	Kozani	13/5/95	N	6.5		6.5				15	15			0.09							149	x	
TR	MT	Dinar	1/10/95	NS	6.2		6.2				10	10										150		
Japa	JP	Kobe	17/1/95	SF		6.9	6.9				30	40			1.9		0.6	1.5					6 x	

#### Description of Table

Column 1, "Loc": Country

Column 2, "REG": Geographic region; NZ=New Zealand, JP=Japan, MT=Eastern Mediterranean, AS=Asia, OC=Oceania, NA=North America, CA=Central America, SA=South America, AU=Australia, AF=Africa

Column 3, "EVENT": Earthquake or fault name

Column 4, "DATE": Date=Earthquake date; paleo = paleoearthquake

Column 5, "ST": Slip type; S=strike-slip, R=reverse-slip, N=normal-slip. Slip types with two letters denote oblique-slip motion, with the order of letters denoting the dominance of one slip type over the other

Column 6, "MsorMjma": Surface wave magnitude

Column 7, "MI": MI=Magnitude derived from interpretation of intensity data

Column 8, "Mw(geol\_subL)": Moment magnitude derived from geological data with the equations  $\log M_0 = 16.1 + 1.5M_w$ , in which  $M_w$ =moment magnitude and  $M_0$ =seismic moment (Column 12)

Column 9, "Mw(geol\_sfcL)": Moment magnitude derived from geological data with the equations  $\log M_0 = 16.1 + 1.5M_w$ , in which  $M_w$ =moment magnitude and  $M_0$ =seismic moment (Column 13)

Column 10, "Mw(seis)": Moment magnitude measured from seismological data

Column 11, "M": The magnitude chosen as the most reliable estimate of the moment magnitude of the earthquake

Column 12, "Mo(geol\_subL)": Seismic moment calculated with the equation  $M_0 = uLW$ , in which  $M_w$ =moment magnitude,  $M_0$ =seismic moment,  $u$ =rigidity modulus ( $3e+11$  dyne/cm<sup>2</sup>),

$L$ =SUBSURFACE rupture length (mean of Columns 14 and 15),  $W$ =rupture width (mean of Columns 18 and 19), and  $D$ =average surface displacement (Column 23)

Column 13, "Mo(geol\_sfcL)": Seismic moment calculated with the equation  $M_0 = uLW$ , in which  $M_w$ =moment magnitude,  $M_0$ =seismic moment,  $u$ =rigidity modulus ( $3e+11$  dyne/cm<sup>2</sup>),  $L$ =SURFACE rupture length (mean of Columns 16 and 17),  $W$ =rupture width (mean of Columns 18 and 19), and  $D$ =average surface displacement (Column 23)

Column 14, "LGTHMsub": Minimum-bound SUBSURFACE rupture length (km)

Column 15, "LGTHMXsub": Maximum-bound SUBSURFACE rupture length (km)

Column 16, "LGTHM": Minimum-bound SURFACE rupture length (km)

Column 17, "LGTHMX": Maximum-bound SURFACE rupture length (km)

Column 18, "WMN": Minimum-bound rupture width (km)

Column 19, "WMX": Maximum-bound rupture width (km)

Column 20, "Dmxsub": Maximum-bound SUBSURFACE displacement (m)

Column 21, "Davsub": Average SUBSURFACE displacement (m)

Column 22, "Dmx": Maximum-bound SURFACE displacement (m)

Column 23, "Dav": Average SURFACE displacement (m)

Column 24, "SRmn": Minimum-bound fault slip rate (mm/yr)

Column 25, "SRmx": Maximum-bound fault slip rate (mm/yr)

Column 26, "SRm": Preferred, or mean fault slip rate (mm/yr)

Column 27, "TSmn": Minimum-bound estimate of total slip (km)

Column 28, "TSM": Maximum-bound estimate of total slip (km)

Column 29, "TSprf": Preferred, mean, or only estimate of total slip (km)

Column 30, "WC": Data are derived from Wells and Coppersmith (1994), and the index number from their original table is shown

Column 31, "AJ": Data are derived from Ambraseys & Jackson (1998), and the index number from their original table is shown

Column 32, "AWS": Data are derived from Anderson et al. (1996), and the index number from their original table is shown.

Column 33, "CS": Data are derived from the PhD work of Christian Stock (pers comm. 1998)

Column 34, "OTHER": Data are derived from the following references; 1=Stirling et al. (1996), 2=Stirling et al. (1998), 3=Cowan et al. (1996), 4=Shimokawa et al. (1998), 5=Tsutsumi et al. (1991), 6=Awata et al. (1995), 7=Kumamoto (1998), 8=Research Group for Active Faults of Japan (1991), 9=Yeats et al. (1997), 10=Mizuno et al. (1991), 11=Russ (1979), 12=Zoback (1979), 13=Barka & Kadinsky-Cade (1988), 14=Peltzer et al. (1988), 15=Ambraseys & Jackson (1990), 16=Hirabayashi et al. (1996), 17=Ritz et al. (1995), 18=Molnar & Deng (1984), 19=Beanland et al. (1989), 20=Dolan et al. (1995), 21=Website of the Southern California Earthquake Center (SCEC), 22=Dokka (1983), 23=Tsukuda et al. (1994), 24=Abercrombie et al. (1998), 25=Yeats et al. (1994), 26=minimum estimate of total slip for dip-slip fault from range-front relief, 26=A.Michetti (pers comm. 1998), 27=Kawasaki (1975), 28=Beanland & Clark (1994)

General Comments: "A" quality events are those for which moment magnitudes are measured directly from seismological data (Column 10), B quality events encompass all the remaining historical earthquakes, and C quality events are those for which moment magnitudes and fault parameters are derived wholly from paleoseismic data. This Table is limited to earthquake magnitudes of  $\geq 6$ .

EOS 2210

INSTITUTE OF GEOLOGICAL AND NUCLEAR SCIENCES LIMITED

Gracefield Research Centre  
69 Gracefield Road  
PO Box 30368  
Lower Hutt  
New Zealand  
Phone +64-4-570 1444  
Fax +64-4-569 0600

Wairakei Research Centre  
State Highway 1, Wairakei  
Private Bag 2000  
Taupo  
New Zealand  
Phone +64-7-374 8211  
Fax +64-7-374 8199

Rafter Laboratory  
30 Gracefield Road  
PO Box 31312  
Lower Hutt  
New Zealand  
Phone +64-4-570 4637  
Fax +64-4-570 4657

Dunedin Research Centre  
764 Cumberland Street  
Private Bag 1930  
Dunedin  
New Zealand  
Phone +64-3-477 4050  
Fax +64-3-477 5232